

## Chapter 4 Soil Physical Properties

### Learning Objectives – Part 1

- Know what color tells you about a soil
- Describe the concept of soil texture and its importance
- Use the textural triangle to determine a soil's textural class based on its sand, silt and clay content

### Soil Physical Properties

- Color
- Texture
- Density (particle density vs. bulk density)
- Pore space (porosity)
- Structure
- Aggregate stability

### Physical properties are important

Physical Properties control plant growth through influence on:

- Soil temperature (darker = warmer)
- Soil aeration (sandy soils well aerated)
- Soil moisture content (clayey soils stay wet)

And they indicate important characteristics of a soil (e.g., lots vs. little organic matter) also... drainage, compaction, consistency, strength, tillth

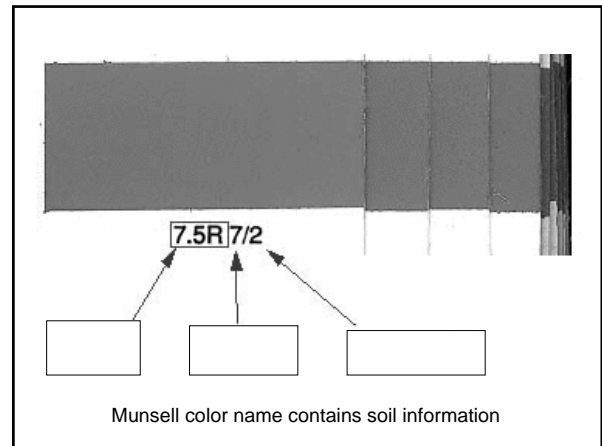
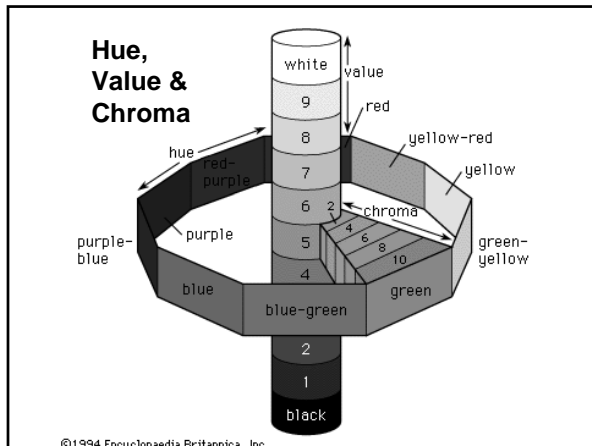
### Physical properties – part 1

- **Color**
  - Caused by coatings on particles:
    - O.M. darkens underlying colors
    - Fe and Al oxides (red & yellow)
    - Moisture (e.g., grey/blue vs. red/brown)
    - Mineralogy (calcite, hematite, manganese, glauconite)
- **Texture**
  - Proportions of sand, silt, and clay
  - Determines water holding capacity, water availability, nutrient supply capacity...

### Color

“quantified” using the Munsell system

- **Hue** (e.g., **5R**) tells you *general shade*; DOES NOT tell you how dark the soil is
- **Value** (e.g., 5R **5/**) tells you *how dark* the soil is: (0 darkest) may indicate current moisture status (dark = wet) and/or amount of organic matter
- **Chroma** (e.g., 5R **5/8**) tells you *color intensity* (0 = gray). Indicator of hydrologic regime (well drained = ↑ O<sub>2</sub> = high chroma)

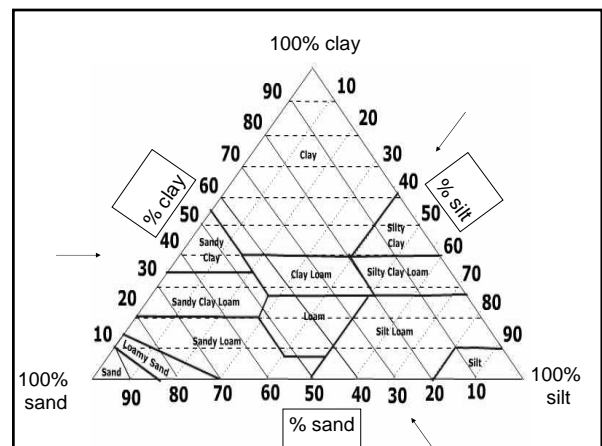
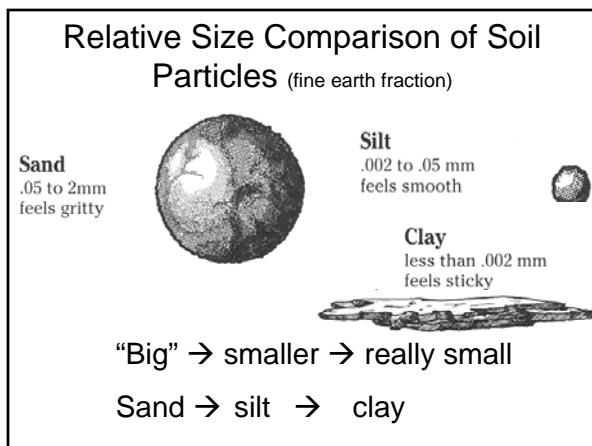


### EXAMPLE TEST QUESTION

- Which soil horizon likely has more organic matter  
**5R** 5/**6**  
 or  
**10YR** 1/**7**?

### Soil Texture

- Proportions of sand, silt, and clay (fine earth fraction)
- Not coarse fragments (>2mm)
- Not OM

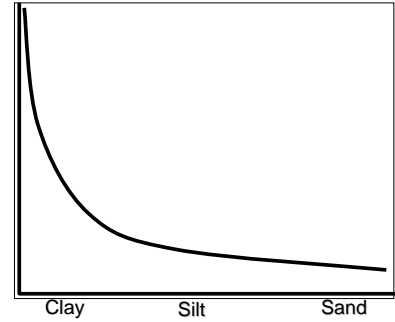


## Texture

- Surface area per unit volume
    - 1 g sand ~ 0.1 m<sup>2</sup> lowest
    - 1 g silt ~ 1 m<sup>2</sup>
    - 1 g clay ~ 10-1000 m<sup>2</sup> highest
  - Large surface area means more charge so greater ability to hold water and nutrients
  - Coarse textured soils have larger pores, fine textured soils have greater total pore space
- ↑ Surface area ↑ weathering

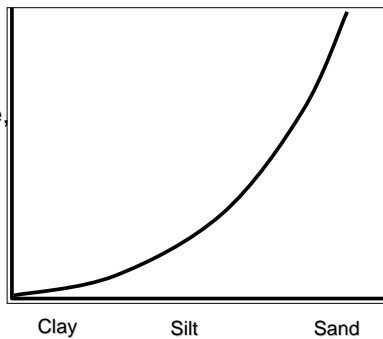
**Gee whiz fact:**  
Fine clay has ~10,000 times as much surface area as the same wt. med. grain sand

Surface area, pore volume, nutrient supply capacity, plasticity and cohesion, swelling



Particle Size

Pore size, infiltration rate, drainage rate, aeration



Particle Size

## Influence of Texture

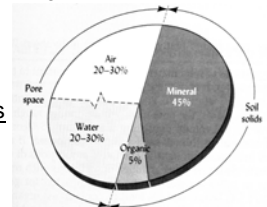
|                        | Sand | Silt   | Clay      |
|------------------------|------|--------|-----------|
| Water-holding capacity | Low  | Medium | High      |
| Aeration               | Good | Medium | Poor      |
| Drainage               | Fast | Slow   | Very slow |
| Nutrient retention     | Low  | Medium | High      |

## Ch. 4 Objectives – Part 2

- Calculate soil bulk density and % pore space
- Discuss factors influencing the above soil properties
- Identify the difference between texture and structure

## Physical properties – part 2

- Density
  - particle density: mass per unit volume (not pore space)  
~ 2.65 Mg/m<sup>3</sup> ( $D_p$  of quartz mineral, dominant mineral in most soils)
  - bulk density: mass per unit volume (including pore space)  
ranges 0.1 – 2.65 Mg/m<sup>3</sup>
- Porosity
  - the volume percentage (%) of the total bulk soil **NOT** occupied by solids



## Bulk Density

Mass of **dry soil** per unit bulk volume,  
INCLUDING pore space

- O.M.  bulk density
- Compaction  bulk density
- Increased bulk density  water infiltration and restricts root growth
- Bulk density is relatively  in coarse-textured soils because

## Typical Bulk Densities

- Histosols (0.1-0.7 Mg/m<sup>3</sup>)
- Cultivated clay and silt loams (0.9-1.5 Mg/m<sup>3</sup>)
- Cultivated sandy loams and sands (1.2-1.8 Mg/m<sup>3</sup>)
- Concrete (ca. 2.4 Mg/m<sup>3</sup>)
- Quartz mineral (2.65 Mg/m<sup>3</sup>)

LOW  
↓  
HIGH

See Fig. 4.12, p. 107 Sulzman

## Typical Bulk Densities: Perhaps worth memorizing...

- OM: ~0.5 Mg/m<sup>3</sup>
- “normal” soils:  $D_b \sim 1.2 - 1.8 \text{ Mg/m}^3$   
– Clayey soils lower than sandy soils
- Upper limit: ~2.65 Mg/m<sup>3</sup>

## Effect of clear-cut on bulk density

Effect varies with soil texture!

## Soil bulk density on skid trails: does tillage help??

## Effect of cultivation on bulk density

|                     | Texture   | Yrs cropped | Increase in $D_b$ due to texture and cultivation |
|---------------------|-----------|-------------|--|
| Coarse<br>↓<br>Fine | Loam      | 58          | 0.18   |
|                     | Silt loam | 40          | 0.26   |
|                     | Clay      | 70          | 0.30   |

**In general, the finer the texture, the more it compacts if aggregates are destroyed**

## Effect of tillage on bulk density

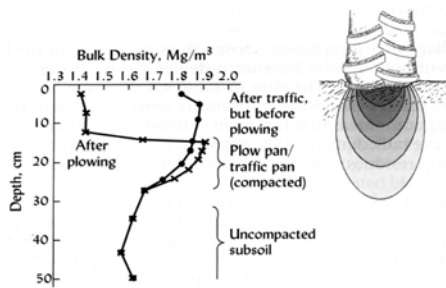
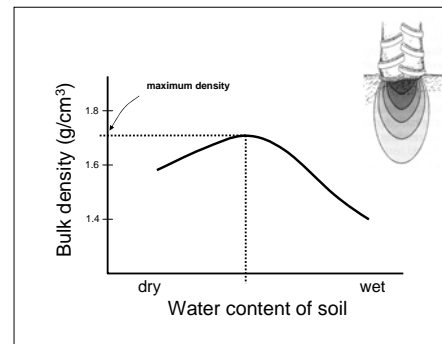


Fig 4.17 p. 113 Sulzman

## Compaction and moisture content



## Effect of recreational use on bulk density

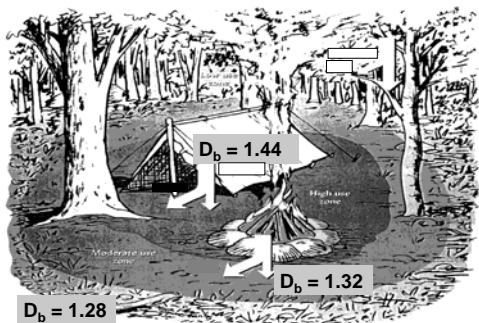
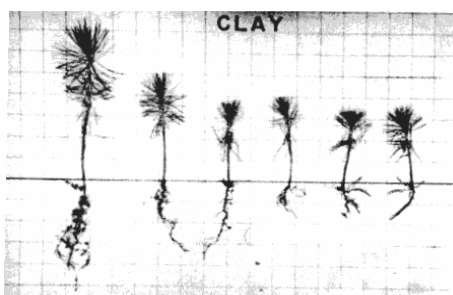


Fig 4.16 p.111 Sulzman

## Effects of compaction (high $D_b$ ) on plant growth

- Silt loam in Arkansas:  
decreasing  $D_b$  by 16% increased cotton yields 13% during a normal year  
59% during a dry year ( $\uparrow$  pore space =  $\uparrow$  water storage capacity)
- Medium textured soils in Washington  
 $D_b > 1.2 \text{ Mg/m}^3$ ; positively correlated with diseases in peas and beans

## Effects of compaction: loblolly pine



Increasing soil compaction  $\rightarrow$  (increasing bulk density)

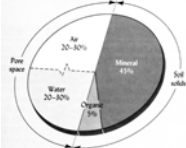
## Again, typical Bulk Densities: Perhaps worth memorizing...

- OM:  $\sim 0.5 \text{ Mg/m}^3$
- "normal" soils:  $D_b \sim 1.2 - 1.8 \text{ Mg/m}^3$   
– Clayey soils lower than sandy soils
- Upper limit:  $\sim 2.65 \text{ Mg/m}^3$

## Particle Density

- Mass of dry soil per unit volume of solids, this EXCLUDES pore space
- So if you compact a soil, does particle density change??

NO!



## Comparison of bulk and particle density

$$g/cm^3 = Mg/m^3$$

In a soil profile, one cubic centimeter (1.0cm<sup>3</sup>) appears like this:  
It contains solids and pore spaces, and the whole cm<sup>3</sup> has a mass of 1.32g

1.32g

If all the solids were compressed to the bottom, the cube would now look like this:

Half contains the pore spaces →

Half contains the solids (Notice the Volume change!)

0.5cm<sup>3</sup>

To calculate Bulk Density of the soil:  
Volume = 1.0cm<sup>3</sup> (Solids and Pores)    Mass = 1.32g (Sieved Solids only)

$$\text{Bulk Density} = \frac{\text{Mass of Dry Soil}}{\text{Volume of soil (Solids and Pores)}}$$

$$\text{Therefore: Bulk Density} = \frac{1.32}{1.0} = 1.32 \text{ g/cm}^3$$

To calculate Particle Density of the soil:  
Volume = 0.5cm<sup>3</sup> (Solids only)    Mass = 1.32g (Sieved Solids only)

$$\text{Particle Density} = \frac{\text{Mass of solids}}{\text{Volume of solids}}$$

$$\text{Therefore: Particle Density} = \frac{1.32}{0.5} = 2.64 \text{ g/cm}^3$$

Similar to Fig 4.10 p. 106 Sulzman

## Pore Space (= porosity)

$$\% \text{ PS} = 100 * (1 - [D_b/D_p])$$

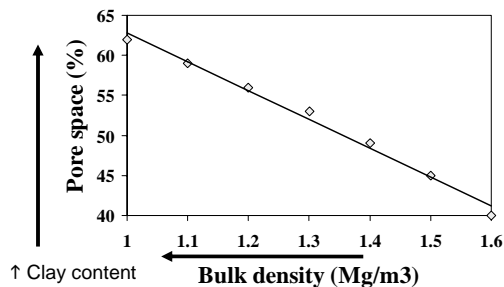
- If you compact or till a soil, the pore space decreases.
- Large pores are filled with air
- Small pores are filled with water
- Clay soils have greater total pore space than sandy soils

## Calculations you must know

$$\text{Bulk Density} = \frac{\text{Mass of Dry Soil}}{\text{Volume of soil (Solids and Pores)}} \quad \text{Particle Density} = \frac{\text{Mass of solids}}{\text{Volume of solids}}$$

$$\% \text{ Pore Space} = 100 * (1 - [D_b/D_p])$$

- A pattern to note: As clay content increases, pore volume increases and bulk density decreases



## Example test question...

- Woodburn silt loam in the OSU quad
- You excavate a hole in the surface horizon and reserve all of the removed material, which you then weigh, dry, and re-weigh

- Freshly collected soil weighs 470 g
- The weight of the dried material is 390 g
- The volume of the hole is 300 cm<sup>3</sup>

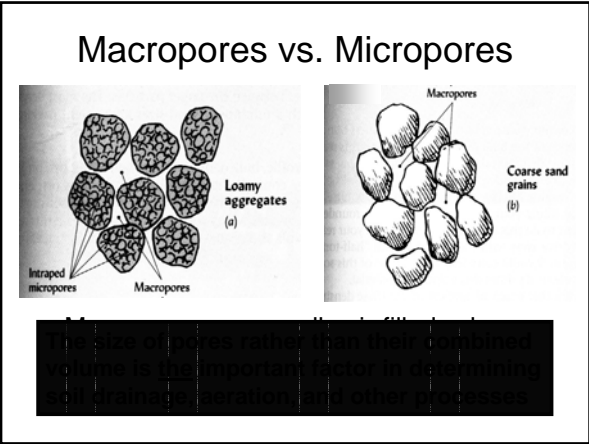
You will recall...

$$D_p = \text{mass dry soil/bulk volume} = 390\text{g}/300\text{cm}^3 = 1.3\text{g/cm}^3$$

$$\% \text{ PS} = 100 * (1 - [D_b/D_p]) = 100(1 - [1.3\text{g/cm}^3 / \text{ }]) = \text{ } \%$$

What is the bulk density of the surface horizon?

What is the % pore space of this horizon?



### Relationship between texture and pore space (data for 10 cm depth)

| Texture          | O.M. (%) | Total pore space (%) | Micropores (%) | Macropores (%) |
|------------------|----------|----------------------|----------------|----------------|
| Sandy loam       | 2        | 42                   | 17             | 25             |
| Silt loam        | 5        | 50                   | 27             | 23             |
| • good structure |          |                      |                |                |
| • poor structure | 5        | 50                   | 40             | 10             |

### Learning Objectives – part 3

- Describe soil structure and its formation and importance
- List factors that promote aggregate stability and state why a gardener or farmer would want a soil with stable aggregates

Structure and aggregate stability related: The same things that lead to strong structure make stable aggregates!

- Structure
    - spatial arrangement of primary soil particles
  - Aggregate stability
    - how easily or not do the peds fall apart?
- Goal of good soil mgmt. { Low bulk density  
Lots of macropores  
Stable aggregates

### Soil Structure

- The arrangement of primary soil particles into groupings called aggregates or peds
- Binding agents provided by plant roots (exudates), organic matter (OM), and clays
- Most important contributor to good structure is OM

### Difference between texture and structure

- Texture = proportions of different particle sizes (% sand, silt, clay)
- Structure = spatial arrangement of those particles

## Why there is structure

- Particles bind because organic compounds and some minerals are “sticky”
  - Biological sources: polysaccharides, proteins, bacterial “glues”
  - Mineralogical sources : oxides, carbonates, silicates (clay particles)

## Characterization of Structure

- *Type*: Shape of the aggregates
  - Four (4) principal shapes
    - Granular (A), platy (E), blocky (B), prismatic (B)
- *Size*
  - fine, medium, coarse
- *Grade*: distinctness (how obvious)
  - Strong, moderate, or weak
- *In general*,
  - if lots of clay → **STRONG** structure, bigger blocks
  - If lots of OM → granular structure

### Structure type: Spheroidal

#### Spheroidal

Characteristic of surface (A) horizons. Subject to wide and rapid changes.

Granular (porous)

Crumb (very porous)

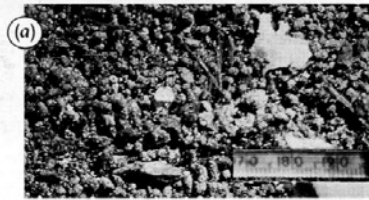
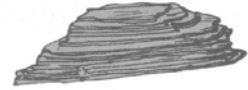


Figure 4.23 p.122 Sulzman

### Structure type: Platey

#### Plate-like

Common in E-horizons, may occur in any part of the profile. Often inherited from parent material of soil, or caused by compaction.



### Structure type: Blocky

#### Block-like

Common in B-horizons, particularly in humid regions. May occur in A-horizons.



Angular blocky



Subangular blocky



### Structure type: Prism-like

#### Prism-like

Usually found in B-horizons. Most common in soils of arid and semi-arid regions.



Columnar (rounded tops)



Sodium induced – disperses clays

Prismatic (flat, angular tops)





## Example of soil structure labeling

GRADE      SIZE      TYPE

- Moderate coarse subangular blocky
- Weak medium platy
- Strong very coarse prismatic
- Moderate very fine granular

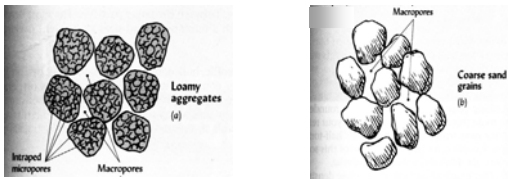
## Aggregates are the units of structure

The same things that lead to strong structure make stable aggregates!

- Amount of OM
- Type and amount of clay
- Amount of stabilizing/flocculating minerals (calcium carbonate, gypsum, etc.)

## Aggregates

- Aggregates contain many small pores; area between aggregates large pores



## Why Are Aggregates Important?

- Increase porosity
- Increase water infiltration, drainage, decrease runoff
- Increase water holding capacity

## When aggregates break down...



## Conditions that Promote Aggregate Stability

1. Low disturbance
2. High root abundance
3. High fungal biomass
4. High OM
5. High clay content

### Effect of OM on aggregate stability

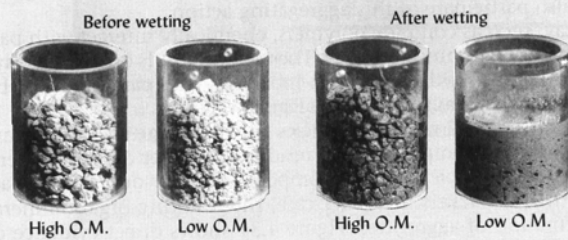


Fig 4.29 p. 128 Sulzman

### Effect of roots & tillage on stability

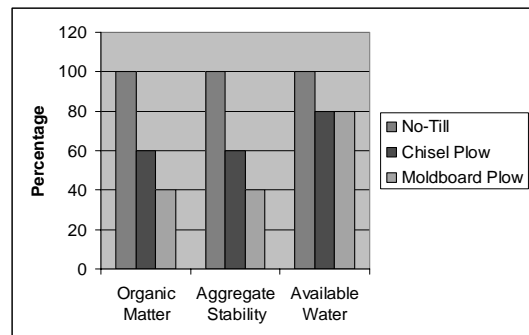


Fig 4.30, p. 129 Sulzman

### Ways to Improve Soil Structure

- Reduce disturbance
- Work soil when dry
- Mulch soil surface
- Add organic materials (crop residues, compost, manure)
- Use cover crops

### Effects of tillage on soil properties



### No-till system – direct seeding



- minimizes the disturbance of soil before planting
- essentially stops soil erosion from wind and water
- saves fuel and time
- greater habitat for birds and other wildlife **BUT** herbicides usually required
- facilitates the buildup of organic matter improving soil quality and yields

### Why would you **NOT** want to use no-till?

- Bird habitat – they eat your seeds
- Residues may contain plant disease
- Weeds!
- Increased structure = better drainage and faster transport of chemicals to the subsurface



### Why we **shouldn't** till wet soils

- Aggregate stability lower when wet
- Once aggregates are gone, pores clog
- Crusts can form, preventing seeds from emerging and water from infiltrating